

Effect of Finite Frequency Bandwidth Limitation on Evaluations of Seismic Radiation Energy of the 1999 Chi-Chi Earthquake

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ABSTRACT

Based on the ω^{-2} and ω^{-3} source models, we explore the effect on estimates of seismic radiation energy, E_s , caused by finite frequency bandwidth limitation of source spectra. Let f_c be the corner frequency of a source spectrum and f_u and f_l are, respectively, the upper and lower bounds of a frequency band in use. Results show that the effect depends on f_u/f_c and f_l/f_c , and E_s is under-estimated when $f_l > 0$ and $f_u < \infty$. When $f_u/f_c < 20$, the effect is sensitive to both f_l/f_c and f_u/f_c for the ω^{-2} source model, but mainly to f_l/f_c for the ω^{-3} model. When $f_u/f_c > 20$, the effect is insensitive only to f_u/f_c for the two models. Let E_s' be the seismic radiation energy estimated without removal of finite frequency bandwidth limitation. Results show: (1) E_s'/E_s first slightly increases and then decreases with increasing f_c ; or (2) E_s'/E_s monotonously decrease with increasing f_c . For the 1999 M_s 7.6 Chi-Chi earthquake, Taiwan, E_s was under-estimated by Hwang et al. (2001), and the degree of under-estimates varies from station to station.

(Key words: Seismic radiation energy, Corner frequency, Finite frequency bandwidth limitation)

1. INTRODUCTION

Seismic radiation energy, E_s , is an important parameter quantifying an earthquake (cf. Wang 2006). However, estimates of E_s can be influenced by the source spectrum, seismic

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radiation patterns, seismic-wave attenuation, surface amplification, site effect, instrumental response, and noise. A correct evaluation of E_s will help seismologists to understand source behavior more exactly. Boore (1988), Di Bona and Rovelli (1988), and Singh and Ordaz (1994) stressed that E_s is underestimated when high-frequency signals are not included. Thus, the E_s measured from local seismograms is usually larger than that done from teleseismic data (Bolt 1986; Smith et al. 1991; Singh and Ordaz 1994; Hwang et al. 2001; Huang et al. 2002). In principle, E_s is measured for $f = 0 - \infty$ Hz, while in practice the measurement can be made only for $f_1 \leq f \leq f_u$ due to limitation in instrumental response and noise. This results in so-called finite frequency bandwidth limitation (denoted *ffbl* hereafter). Ide and Beroza (2001) theoretically studied such an effect in a high-frequency regime. Based on the ω^{-2} source model, Wang (2004) studied the effect in both low- and high-frequency regimes. Both studies show underestimation of E_s due to the *ffbl*-effect.

For the 1999 M_s 7.6 Chi-Chi earthquake, Wang (2004) made corrections only based on the ω^{-2} source model. However, Huang and Wang (2002) stressed that a ω^{-3} model must be taken into account for the northern fault plane of the earthquake. In this work, the *ffbl*-effects of source spectrum for both high- and low-frequency regimes on estimates of E_s based on the ω^{-2} and ω^{-3} source models (Aki 1967; Brune 1970) will be discussed in detail. Theoretical results will be applied to correct estimates of the seismic radiation energy of the Chi-Chi earthquake.

2. DEFINITION AND METHODOLOGY FOR MEASURING E_s

The source spectra of earthquakes are mainly controlled by the low-frequency spectral level (Ω_0) and corner frequency (f_c) (Aki 1967). Theory and observations show that when $f > f_c$, the spectral amplitude decays in a power-law function like $f^{-\alpha}$. Commonly accepted power-law functions have either f^{-2} or f^{-3} , which are, respectively, referred to as the ω^{-2} and ω^{-3} source models, where $\omega = 2\pi f$ (Aki 1967; Brune 1970).

Let $d(t)$ and $v(t)$ be the source displacement and velocity, respectively. Their Fourier transforms are, respectively, $D(f)$ and $V(f)$. $D(f)$ can be approximated by $D_2(f) = \Omega_0 / [1 + (f/f_c)^2]$ for the ω^{-2} model and $D_3(f) = \Omega_0 / [1 + (f/f_c)^2]^{3/2}$ for the ω^{-3} one (cf. Beresnev and Atkinson 1997). Hence, the approximations of $V(f)$ are, respectively:

$$V_2(f) = 2\pi f \Omega_0 / [1 + (f/f_c)^2] \quad , \quad (1)$$

$$V_3(f) = 2\pi f \Omega_0 / [1 + (f/f_c)^2]^{3/2} \quad . \quad (2)$$

As $f \ll f_c$, $V_2(f) \sim f^1$ and $V_3(f) \sim f^1$, while as $f \gg f_c$, $V_2(f) \sim f^{-1}$ and $V_3(f) \sim f^{-2}$. Eqs. (1) ~ (2) can be approximated individually by a piece-wise linear function (Fig. 1).

E_s is calculated by the following expression:

$$E_s = 4\pi\rho\beta \int v^2(t)dt = 4\pi\rho\beta [2 \int V^2(f)df] \quad , \quad (3)$$

where ρ and β are, respectively, the density and the S-wave velocity. In principle, the first integral is performed from $-\infty$ to $+\infty$ in the time domain and the second one from 0 to $+\infty$ in the frequency domain. Define:

$$I_V = \int v^2(t)dt = 2 \int V^2(f)df \quad (4)$$

Thus, $E_s = 4\pi\rho\beta I_V$.

3. THE EFFECT DUE TO FINITE FREQUENCY BANDWIDTH LIMITATION

Wang (2004) derived the formulas to show the *ffbl*-effect based on the ω^{-2} model. For the purpose of comparison, his formulas are shown again below. In the following, we add a subscript 'o' to denote a quantity obtained through integration from $-\infty$ and $+\infty$ sec in the time domain or from 0 to ∞ Hz in the frequency domain. Inserting Eqs. (1) and (2) into Eq. (4), respectively, leads to:

$$I_{V_{2o}} = I_V(f_l = 0, f_u = \infty) = \Omega_o^2 (2\pi f_c)^3 / 4 \quad (5)$$

$$I_{V_{3o}} = I_V(f_l = 0, f_u = \infty) = \Omega_o^2 (2\pi f_c)^3 / 16 \quad (6)$$

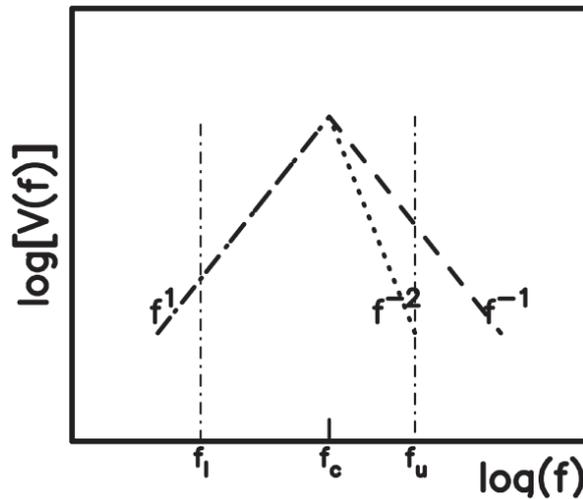


Fig. 1. The log-log plots of the normalized, simplified velocity spectra, $V(f)$ versus frequency, f : the dashed and dotted lines, respectively, for the f^{-1} and f^{-2} source velocity models. The two vertical dashed-dotted lines display the frequency band in use.

where the subscript is 2 for the ω^{-2} model and 3 for the ω^{-3} model. Clearly, $I_{V20} = 4I_{V30}$. When integration is made only in a finite frequency band from f_l to f_u , with $f_l < f_c < f_u$, which is in between two dashed-dotted lines as shown in Fig. 1, the *ffbl*-effect exists. When $f_u/f_l = f_u/f_c$, for the ω^{-2} model the high-frequency cut-off part with $f > f_u$ is almost equal to that from the low-frequency one with $f < f_l$; while for the ω^{-3} model the former is smaller than the latter.

Inserting Eqs. (1) and (2) into Eq. (4), with $f_l < f_c < f_u$, respectively, gives:

$$I_{V2} = 2\Omega_0^2 \int (2\pi f)^2 [1 + (f/f_c)^2]^{-2} df \quad , \quad (7)$$

$$I_{V3} = 2\Omega_0^2 \int (2\pi f)^2 [1 + (f/f_c)^2]^{-3} df \quad , \quad (8)$$

where the integral range is of from f_l to f_u . After integration, Eqs. (7) and (8), respectively, becomes:

$$I_{V2} = I_{V20} F_{V2} \quad , \quad (9)$$

$$I_{V3} = I_{V30} F_{V3} \quad , \quad (10)$$

where

$$F_{V2} = (2/\pi) \{ -(f_u/f_c) / [1 + (f_u/f_c)^2] + (f_l/f_c) / [1 + (f_l/f_c)^2] + \tan^{-1}(f_u/f_c) - \tan^{-1}(f_l/f_c) \} \quad , \quad (11)$$

$$F_{V3} = (4/\pi) \{ -(f_u/f_c) / [1 + (f_u/f_c)^2]^2 + (f_u/f_c) / 2 [1 + (f_u/f_c)^2] + \tan^{-1}(f_u/f_c) / 2 + (f_l/f_c) / [1 + (f_l/f_c)^2]^2 - (f_l/f_c) / 2 [1 + (f_l/f_c)^2] - \tan^{-1}(f_l/f_c) / 2 \} \quad . \quad (12)$$

When $f_l = 0$ and $f_u \rightarrow \infty$, $F_{V2} = 1$ and $F_{V3} = 1$, and, thus, $I_{V2} = I_{V20}$ and $I_{V3} = I_{V30}$.

Hereafter, let E_s and E_s' be the values of seismic radiation energy estimated, respectively, with and without removal of the *ffbl*-effect. From Eqs. (9) - (12), the energy ratio of E_s' to E_s is:

$$E_{s2}'/E_s = F_{V2} \quad , \quad (13)$$

for the ω^{-2} model and:

$$E_{s3}'/E_s = F_{V3} \quad , \quad (14)$$

for the ω^{-3} model. The variations of E_{s2}'/E_s and E_{s3}'/E_s with f_l/f_c are made only for $f_l/f_c < 1$ and $f_l/f_c > 1$ under the request of $f_l < f_c < f_u$. In other words, the calculations are made when $f_l/f_c = 0.05 - 0.95$ and $f_u/f_c = 2$ to 20 , with a difference of 2 . The plots for ten values of f_u/f_c are shown, respectively, in Fig. 2 (for E_{s2}'/E_s) and Fig. 3 (for E_{s3}'/E_s), where the dotted line displays the energy ratio of 1 , without *ffbl*.

In Figs. 2 and 3, all curves are below the dotted line with $E_s'/E_s = 1$, and, thus, E_{s2}'/E_s and E_{s3}'/E_s are both smaller than 1 , with a maximum of about 0.937 for E_{s2}'/E_s and 0.999 for E_{s3}'/E_s . Obviously, the *ffbl*-effect yields an under-estimation of seismic radiation energy. E_{s2}'/E_s and

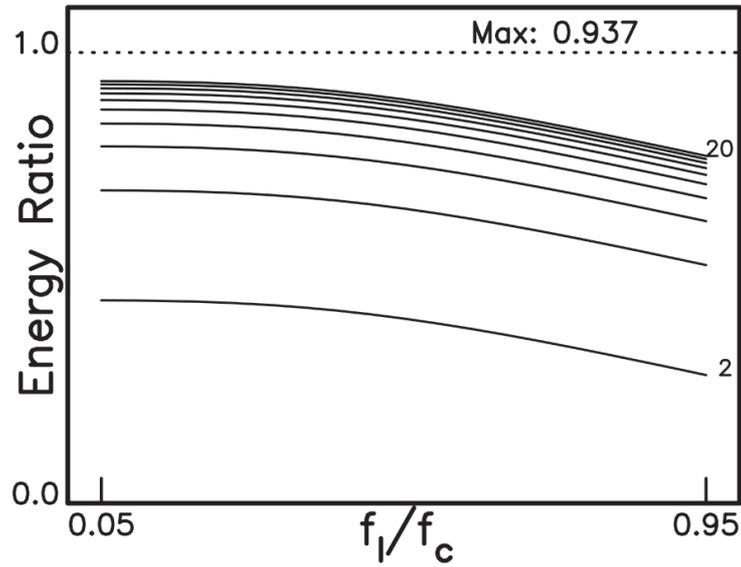


Fig. 2. The variations of E_{s2}'/E_s with f_l/f_c (from 0.05 to 0.95) for ten values of f_0/f_c (from 2 to 20). The dotted line represents $E_{s2}'/E_s = 1$.

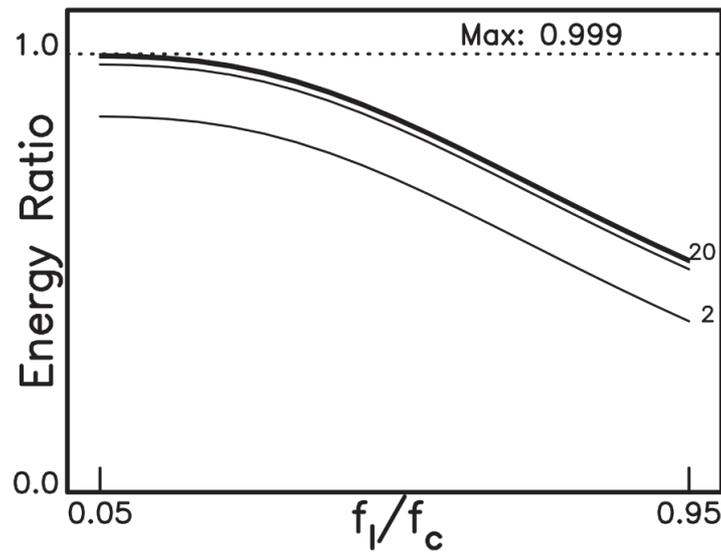


Fig. 3. The variations of E_{s3}'/E_s with f_l/f_c (from 0.05 to 0.95) for ten values of f_0/f_c (from 2 to 20). The dotted line represents $E_{s3}'/E_s = 1$.

E_{s3}'/E_s both decrease with increasing f_l/f_c , and the amount of the decreasing rate increases with f_l/f_c . For fixed f_c , decreases in E_{s2}'/E_s and E_{s3}'/E_s with increasing f_l/f_c lead to increases in E_{s2}'/E_s and E_{s3}'/E_s with decreasing f_l . This indicates that an increase in the width of the low-frequency regime improves estimation of E_s . When $f_l/f_c < 0.4$ for E_{s2}'/E_s and $f_l/f_c < 0.2$ for E_{s3}'/E_s , the curves are almost flat for all f_u/f_c . This means that $f_l = 0.4f_c$ for E_{s2}'/E_s and $f_l = 0.2f_c$ for E_{s3}'/E_s are the individual optimum lower bounds to lead to a stable value of E_s .

E_{s2}'/E_s and E_{s3}'/E_s both increase with f_u/f_c . The curves are close to one another for E_{s2}'/E_s when $f_u/f_c \geq 10$ and for E_{s3}'/E_s when $f_u/f_c \geq 4$, thus indicating that $f_u = 10f_c$ for E_{s2}'/E_s and $f_u = 4f_c$ for E_{s3}'/E_s are both large enough to lead to a stable estimate of E_s . For fixed f_c , increases in E_{s2}'/E_s and E_{s3}'/E_s with f_u/f_c yield increases in E_{s2}'/E_s and E_{s3}'/E_s with f_u , thus indicating that an increase in the width of high-frequency regime improves estimates of E_s . This is consistent with others' (Boore 1988; Di Bona and Rovelli 1988; Singh and Ordaz 1994; Ide and Beroza 2001).

Figures. 2 and 3 show that for fixed f_l , decreases in E_{s2}'/E_s and E_{s3}'/E_s with increasing f_l/f_c lead to increases in E_{s2}'/E_s and E_{s3}'/E_s with f_c , thus implying that the *ffbl*-effect in the low-frequency regime gives a greater underestimate of E_s for events with lower f_c than for those with higher f_c . This effect is stronger for the ω^{-3} model than the ω^{-2} model. For fixed f_u , increases in E_{s2}'/E_s and E_{s3}'/E_s with f_u/f_c result in increases in E_{s2}'/E_s and E_{s3}'/E_s with decreasing f_c , thus showing that the *ffbl*-effect in the high-frequency regime yields a bigger underestimate of E_s for events with higher f_c than for those with lower f_c . When both f_l and f_u are finite and fixed, an increase in f_c will lead to a decrease in both f_l/f_c and f_u/f_c . Hence, the variation of E_{s2}'/E_s and E_{s3}'/E_s with f_c can be either of the following two types: (1) the ratio first slightly increases and then decreases with increasing f_c ; and (2) the ratio monotonously decrease with increasing f_c .

4. RE-EVALUATION OF E_s OF THE 1999 CHI-CHI EARTHQUAKE

The M_s 7.6 Chi-Chi earthquake, which ruptured the Chelungpu fault, struck central Taiwan on 20 September 1999. The epicenter and the fault trace are displayed in Fig. 4. The values of f_c and Ω_0 at four near-fault stations evaluated by Hwang et al. (2001) are $f_c = 0.064 - 0.193$ Hz and $\Omega_0 = 89.4 - 2350.0$ cm. The values of f_c and Ω_0 are shown in columns 2 and 3 of Table 1. They also estimated the values of E_s , which is equivalent to E_{s2}' for the ω^{-2} model and E_{s3}' for the ω^{-3} source model in this study and denoted by E_s' in column 8 of Table 1, at four near-fault seismic stations (see Fig. 4) based on two sets of f_l and f_u : (1) $f_l = 0.03$ and $f_u = 1.0$ Hz at TCU102 and TCU052; and (2) $f_l = 0.03$ and $f_u = 3.0$ Hz at TCU076 and TCU129. The values of f_l and f_u used are shown in columns 4 and 5 of Table 1. In order to obtain a reliable value of E_s , they eliminated the effects caused by seismic radiation patterns, seismic-wave attenuation, surface amplification, site effect, and instrumental response. Wang (2004) re-evaluated the values of E_s estimated by Hwang et al. (2001) through the removal of the *ffbl*-effect based on the ω^{-2} model. His values of E_{s2}'/E_s and E_s are shown in parentheses of columns 9 and 10 in Table 1.

From the values of f_c , f_l , and f_u at the four stations, the ratios of f_l/f_c and f_u/f_c are calculated and given in column 6 and 7 of Table 1: f_l/f_c of from 0.155 to 0.469 and f_u/f_c of from 8.197 to

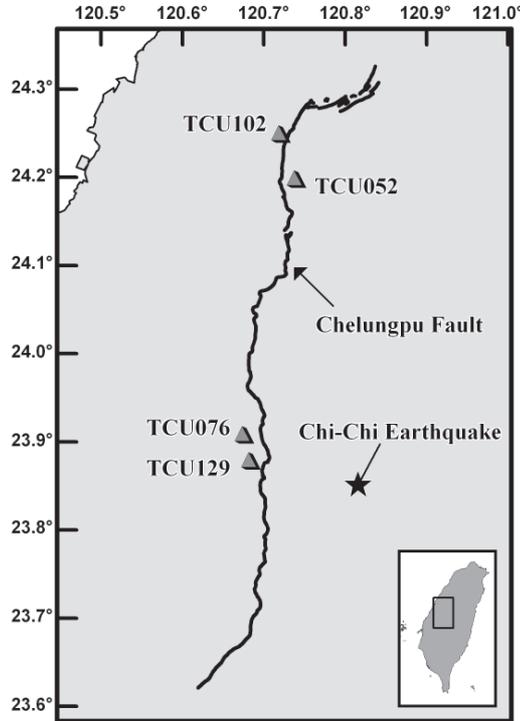


Fig. 4. A map showing the epicenter (in a solid star) of the 1999 Chi-Chi earthquake, the Chelungpu fault (in a solid line), and four near-fault seismic stations (in solid triangles).

Table 1. The values of several parameters at four near-fault seismic stations. In columns 9 and 10, E_s'/E_s and E_s , respectively, includes E_{s2}'/E_s and E_s for the ω^{-2} model and E_{s3}'/E_s and E_s for the ω^{-3} model. The values of E_s'/E_s and E_s not inside the parenthesis are, respectively, E_{s2}'/E_s and E_s taken from Wang (2004). The values of E_s'/E_s and E_s inside the parenthesis are, respectively, E_{s3}'/E_s and E_s of this study.

Stations	f_c (Hz)	Ω_0 (cm)	f_l (Hz)	f_u (Hz)	f_l/f_c	f_u/f_c	E_s' (ergs)	E_s'/E_s	E_s (erg)
TCU102	0.122	551.2	0.03	1.0	0.246	8.197	9.7×10^{22}	0.840 (0.974)	1.2×10^{23} (9.9×10^{22})
TCU052	0.064	2350.0	0.03	1.0	0.469	15.625	2.5×10^{23}	0.884 (0.877)	2.8×10^{23} (2.9×10^{23})
TCU076	0.193	89.4	0.03	3.0	0.155	15.544	1.0×10^{22}	0.917	1.1×10^{22}
TCU129	0.160	105.0	0.03	3.0	0.188	18.750	6.7×10^{21}	0.929	7.2×10^{21}

18.750. The values of E_{s2}'/E_s and E_s re-evaluated by Wang (2004) based on the ω^{-3} model are shown in the parentheses of columns 9 and 10 in Table 1. Clearly, the *ffbl*-effect results in an underestimate of E_s , and the underestimate is higher at two northern stations than at the southern ones. We calculate the values of E_{s2}'/E_s and E_s at two northern stations using Eqs. (12) and (13) based on the ω^{-3} model. Results are shown in the parentheses of columns 9 and 10 of Table 1. Obviously, the results are opposite to those evaluated based on the ω^{-2} model. The difference is bigger at TCU102 and smaller at TCU052. Based on the ω^{-3} model, the value of E_s at TCU102 estimated by Hwang et al. (2001) is good enough.

To examine the problem in advance, we plot the variations of energy ratio with f_c in the range 0.05 - 0.20 Hz in Fig. 5 for two sets of f_1 and f_u : (1) $f_1 = 0.03$ and $f_u = 1.0$ Hz for northern seismic stations; and (2) $f_1 = 0.03$ and $f_u = 3.0$ Hz for southern ones. The dashed and solid lines represent, respectively, E_{s2}'/E_s and E_{s3}'/E_s for the northern stations; and the dashed-dotted line shows E_{s2}'/E_s for the southern stations. The estimated results are also plotted by an open circle or a cross attached with a station code in Fig. 5.

In Fig. 5, E_{s2}'/E_s first increases and then decreases with increasing f_c . Whereas, E_{s3}'/E_s first increases with f_c and then becomes flat when $f_c > 0.14$ Hz. The variations are as expected as mentioned above. The three variations are all below the dotted line with $E_s'/E_s = 1$, thus showing underestimate of E_s at the four seismic stations. The solid line intersects the dashed

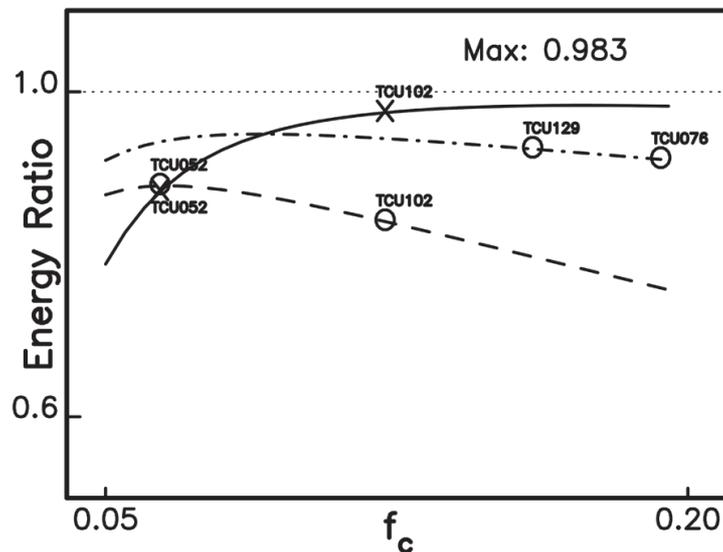


Fig. 5. The variations of energy ratio with f_c for various values of f_1 and f_u as mentioned in the text: the dashed and solid lines, respectively, for E_{s2}'/E_s and E_{s3}'/E_s at the northern stations, and the dashed-dotted line for E_{s2}'/E_s at the southern ones. The related values at four near-fault seismic stations for the Chi-Chi earthquake are displayed by an open circle or a cross attached with a station code. The dotted line represents $E_s'/E_s = 1$.

and dashed-dotted ones at $f_c = 0.065$ and $f_c = 0.090$ Hz, respectively. Hence, at the northern stations the underestimate of E_s is smaller from the ω^{-2} model than from the ω^{-3} model when $f_c < 0.065$ Hz, and opposite when $f_c > 0.065$ Hz. The difference between the effects from the two models is small at TCU052 and large at TCU102. Underestimation of E_s is smaller at the southern stations than at the northern ones when $f_c < 0.09$ Hz, and opposite when $f_c > 0.09$ Hz. Consequently, the values of E_s at the four near-fields suggested by this study are 9.9×10^{22} erg at TCU102, 2.9×10^{23} erg at TCU052, 1.1×10^{22} erg at TCU076, and 7.2×10^{21} erg at TCU129.

5. CONCLUSION

The *ffbl*-effect of source spectrum on estimation of seismic radiation energy, E_s , is analyzed theoretically on the basis of the ω^{-2} and ω^{-3} source models. Such an effect depends on f_l/f_c and f_u/f_c . Numerical results obviously show that E_s are underestimated for all f_l/f_c and f_u/f_c . An increase in the frequency bandwidth including either the high- or low-frequency regime will increase reliability of estimating E_s . When $f_u/f_c < 20$, the effect is sensitive to both f_l/f_c and f_u/f_c for the ω^{-2} model, but mainly to f_l/f_c for the ω^{-3} model. When $f_u/f_c > 20$, the effect is insensitive to f_u/f_c for the two models. For the two source models, E_s'/E_s depends on f_c in either: (1) E_s'/E_s first slightly increases and then decreases with increasing f_c ; or (2) E_s'/E_s monotonously decreases with increasing f_c . Numerical results also suggest that Fig. 2 or 3 together with Fig. 5 can help us to select an appropriate frequency band for estimating a reliable value of seismic radiation energy. The values of f_l and f_u leading to an optimum estimate of E_s are $f_l = 0.4f_c$ and $f_u = 10f_c$ for the ω^{-2} model and $f_l = 0.2f_c$ and $f_u = 4f_c$ the ω^{-3} model.

For the 1999 M_s 7.6 Chi-Chi, Taiwan, earthquake, the revised values of E_s show that E_s was underestimated by Hwang et al. (2001). However, the degree of underestimates varies from station to station. At the northern stations underestimation of E_s is smaller for the ω^{-2} model than for the ω^{-3} model when $f_c < 0.065$ Hz, and opposite when $f_c > 0.065$ Hz. The difference between the effects from the two models is small at TCU052 and large at TCU102. Underestimation of E_s is smaller at the southern stations than at the northern ones when $f_c < 0.09$ Hz, and opposite when $f_c > 0.09$ Hz. The values of E_s at the four near-field are 9.9×10^{22} erg at TCU102, 2.9×10^{23} erg at TCU052, 1.1×10^{22} erg at TCU076, and 7.2×10^{21} erg at TCU129.

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